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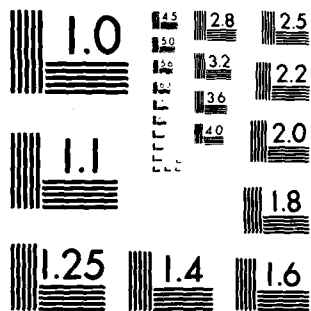
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Evaluation of Geothermal Potential of Range Bravo 17 and the Shoal Site, Naval Air Station, Fallon

by

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Public Works Department

MARCH 1980

**NAVAL WEAPONS CENTER
CHINA LAKE, CALIFORNIA 93555**



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Naval Weapons Center

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FOREWORD

The Naval Weapons Center has been tasked with evaluating the geothermal potential of naval installations throughout the world. The evaluation of the Shoal Site and Range Bravo 17 was conducted by the Naval Weapons Center during the period October 1978 through September 1979, under sponsorship of the Naval Civil Engineering Laboratory, Project No. Z0840-SL.

This report has been reviewed for technical accuracy by Dr. Carl F. Austin, Head, Utilization Division, Public Works Department.

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1 March 1980

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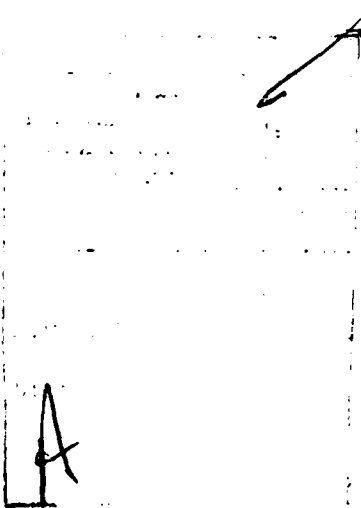
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(U) *Evaluation of Geothermal Potential of Range Bravo 17 and the Shoal Site, Naval Air Station, Fallon* (U), by J. A. Whelan, C. R. Rodgers, J. Bown, and Jack Neffew. China Lake, Calif., Naval Weapons Center, March 1980. 34 pp. (NWC TP 6142, publication UNCLASSIFIED.)

(U) As part of the task of evaluating the geothermal potential of naval installations throughout the world, Range Bravo 17 and the Shoal Site, both located at the Naval Air Station, Fallon, Nev., were evaluated.

(U) An aeromagnetic survey was conducted; seismic refraction studies and extensive hydrologic studies were made; thermal gradient and heat flow data were taken; and water chemistry and mercury geochemistry analyses were made.

(U) The results of the study show that the geothermal potential of the Shoal area and Range Bravo 17 is low. It was concluded that further exploration of these sites is not warranted at this time.



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INTRODUCTION

It is important for the Navy to know the geothermal potential of its lands (1) to prevent or minimize encroachment, and (2) if the resources warrant, to develop energy self-sufficiency. Consequently, the Geothermal Utilization Division, Public Works Department, Naval Weapons Center (NWC), has been tasked with evaluating the geothermal potential of naval installations throughout the world.

As part of this task, the geothermal potential of Range Bravo 17 and the Shoal Site, both located at the Naval Air Station (NAS), Fallon, Nevada, was evaluated. The studies consisted of a literature review, a reconnaissance of the Shoal Site, and soil sampling for mercury over much of Range Bravo 17. Results of these studies are given in this report. Because of the large amount of data available in the literature (particularly in two reports on the Shoal Site^{1, 2}), the limited field work undertaken was considered adequate for an evaluation of the geothermal potential of the area. Additional work is not recommended.

GEOGRAPHY

The Naval Air Station, Fallon, and its associated ranges are located in a portion of Nevada with generally high geothermal potential. Stillwater-Soda Lake, Desert Peak, and Brady-Hazen, designated Known Geothermal Resource Areas (KGRA), respectively lie to the northwest and northeast of NAS Fallon, with the Stillwater-Soda Lake KGRA bordering the station to the north. Just southeast of NAS Fallon is the Salt Wells Basin KGRA. Lees Hot Springs is on the northeast corner of Range Bravo 19, and Oxygeothermal has drilled an observation hole there. The Electronic Warfare Range is in Dixie Valley. To the north are Dixie Hot Springs and the Dixie Valley KGRA. Sun Oil Company has drilled a commercial well in Dixie Valley. Range Bravo 17 is in Fairview Valley, a southerly extension of Dixie Valley. Range Bravo 20 has the Stillwater-Soda Lake KGRA to the south and the Brady-Hazen KGRA to the west.

Range Bravo 17, one of the areas evaluated in this study, lies to the south of U.S. Highway 50, about 40 miles southwest of Fallon and 75 miles east of Austin, Nevada. The Electronic Warfare Training Area lies just to the north of Highway 50 in the same area. Range Bravo 17 occupies a large portion of T16N, R33E, and about the western third of T16N, R34E (MDB and M).

The results of evaluating three additional areas are also presented in this report. These are two target areas—the northernmost consisting of S 1/2 Sec 22, S 1/2 Sec 32, SW 1/2

¹Nevada Bureau of Mines. *Geological, Geophysical, and Hydrological Investigation of the Sand Springs Range, Fairview Valley and Fourmile, Churchill County, Nevada*. Reno, Nev., University of Nevada, 1962. 128 pp. (Publication UNCLASSIFIED.)

²———. *Project 33.2 Geological, Geophysical, Chemical, and Hydrological Investigations of the Sand Springs Range, Fairview Valley, and Fourmile Flat, Churchill County, Nevada*. Reno, Nev., University of Nevada, October 1963. 360 pp. (VUF-1001, publication UNCLASSIFIED.)

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Sec 24, W 1/2 Sec 25, Sec 26, and Sec 27, T16N, R32E (MDB and M); and the southerly one consisting of Sec 10, Sec 11, W 1/2 Sec 12, NW 1/4 Sec 13, N 1/2 Sec 14, and N 1/2 Sec 15, T16N, R32E (MDB and M)—and the lands between these two. In this latter area, which has been designated the Shoal Site, a 12.4-kiloton atomic device was detonated at a depth of 1,205 feet on 26 October 1963. Access to the target areas and Shoal Site is by U.S. Highway 50, south on Nevada State Highway 31 for 6.1 miles, and east 4 miles into the Sand Springs Range. Locations are given in Figure 1, and an air-photo of Fairview Valley is given as Figure 2. (No drilling or excavating is allowed within 3,500 feet of ground zero.)

CLIMATE

The Sand Springs Range and Fairview Valley have a typical Great Basin arid climate with about 5 inches of rain in the valley and 12 inches at the highest elevation (footnote 1).

The mean annual temperature at Fallon, as reported in a 1961 annual climate summary, is 50.6°F. January is the coldest month and July the warmest. The mean January temperature is about 33°F and that for July about 76°F. Extremes for those months are -14°F and 105°F (see Appendix A).

GEOLOGY

Both before and after the atomic test of 1963, detailed geological, geophysical, and hydrologic studies of the Shoal Site area were made by the Nevada Bureau of Mines, the Nevada Mining Analytical Laboratory, and the Desert Research Institute, all of the University of Nevada. Copies of the two reports prepared under the test contract (footnotes 1 and 2) were obtained for the writers by Robert Clarke, Nevada Operations, Department of Energy. Geologic data from these reports were used extensively in the present study.

A good general summary of the geology of the Sand Springs Range and Fairview Valley is given in the Nevada Bureau of Mines' first Shoal report (footnote 1); the following description is quoted from that report.

"The Sand Springs Range is made up chiefly of a Cretaceous granite intrusive body bordered on both the north and south by Mesozoic (?) metamorphic rocks.... Locally both the granite body and metamorphic rocks are overlain by Tertiary and Quaternary volcanic rocks.... Numerous aplite-pegmatite dikes cut the granitic body; most of these dikes are concentrated in a zone extending south along the western crest of the Range and swinging southeast to the granite-metamorphic contact. Many andesite and rhyolite dikes intrude the granitic body, the metamorphic rocks, and the aplite-pegmatite dikes. Tertiary and Quaternary alluvial and eolian deposits occupy the valleys to the east and west of the Range.

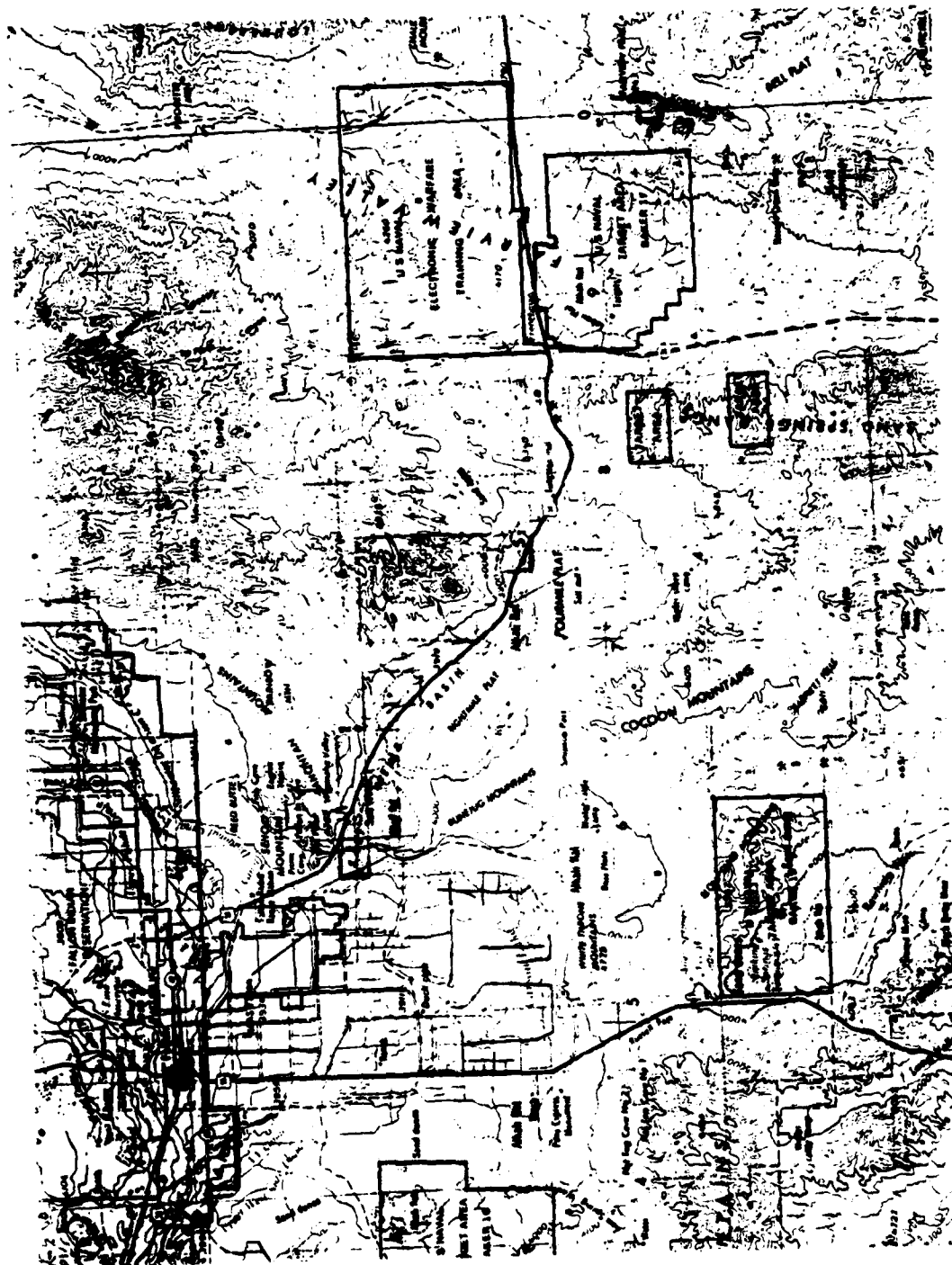


FIGURE 1. Location of Shoal Site and Range Bravo 17.

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FIGURE 2. Aerial Photograph of Fairview Valley.

"Although the Range is a north-south trending fault block, north-south faults are rare, the Range having been uplifted along a series of northwest-and-northeast-trending high-angle faults, which form a sawtoothed pattern in plan. The down-dropped Fairview Valley block to the east contains over 5,000 feet of unconsolidated sediments; in contrast, the Fourmile Flat area to the west is a pediment, thinly veneered with alluvium near the Range, the alluvium thickening to about 1,300 feet immediately south of the salt flat.

"The structural pattern is remarkably consistent throughout much of the Range. A system of steeply-dipping faults and joints trending about N 60° W cuts the granitic body; most of the dikes are intruded along these structures. A second system of steeply-dipping faults and parallel, closely-spaced fracture cleavage, trending more or less N 30° E, cuts the granitic body and dikes. Other directions of faulting, jointing, and cleavage are common locally. A gently-north-dipping thrust cuts the metamorphic rocks just south of U.S. Highway 50; there is no evidence of thrusting elsewhere in the Range."

A geological map, adapted from the Nevada Bureau of Mines' second Shoal report (footnote 2) is given as Figure 3.

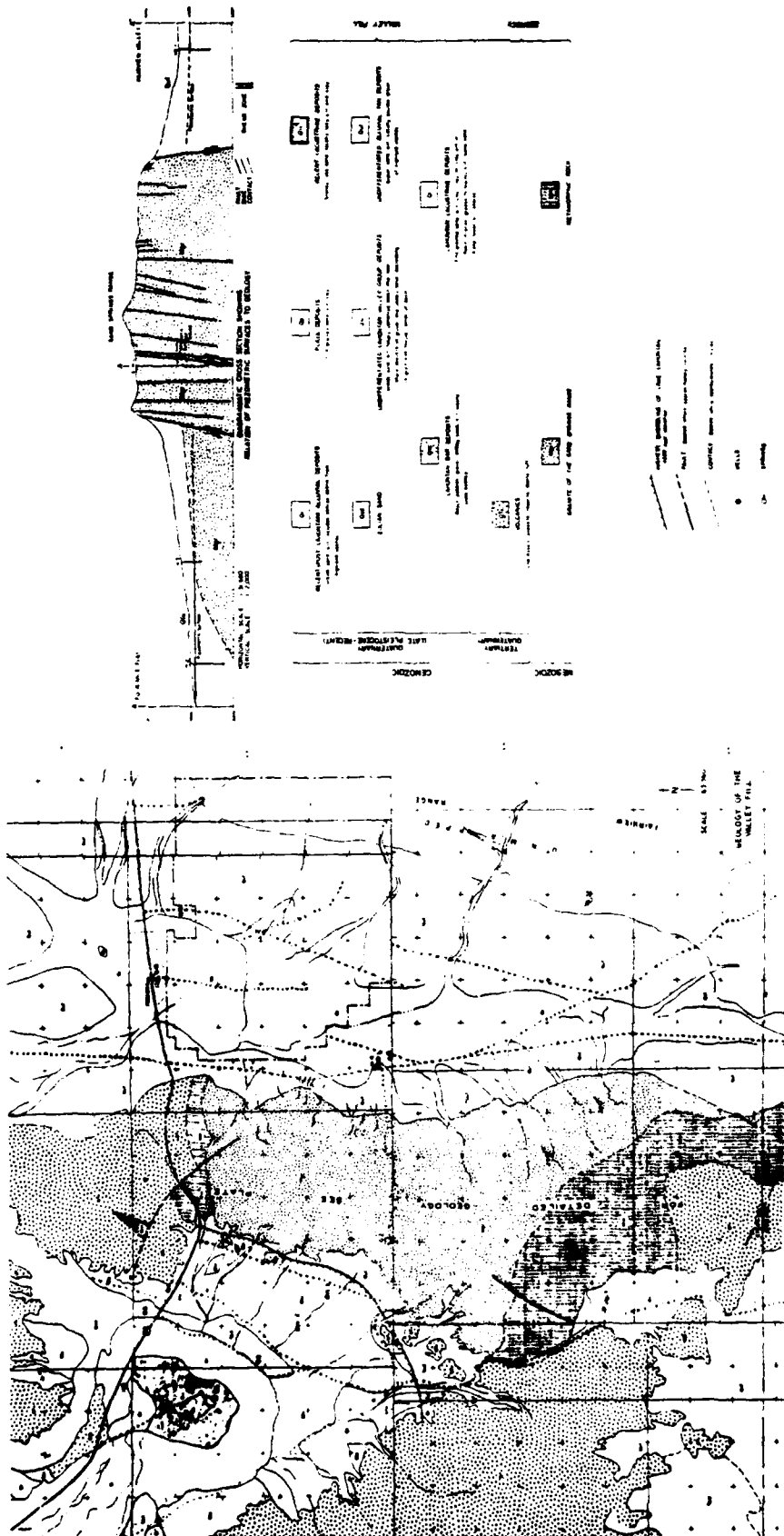
The undesignated target areas and Shoal Site are in the granitic portion of the range. The intrusive rock varies from porphyritic biotite granite to granodiorite, with the granodiorite predominating in the west central part of the range (footnote 1).

The granite is a porphyritic biotite granite with abundant, large orthoclase phenocrysts in a medium-to-coarse grained groundmass of quartz, orthoclase, plagioclase, varying amounts of small biotite flakes and/or books, and locally some hornblende.

The granodiorite appears to have both intrusive and gradational relationships with the granite. Potassium-argon age determination on biotite from the granite and granodiorite was 79.6 (± 2.0) and 76.0 (± 2.0) million years respectively.

The aplite pegmatite dikes extend from less than 1 inch to more than 20 feet in width. Aplitic and pegmatitic material commonly occur in the same dike, usually with sharp contacts. In most cases, the aplite forms the borders of the dikes with pegmatite occurring as lenses, stringers, and continuous layers in the centers. Some dikes change to quartz veins along their strike. The aplite consists of 30 to 40% quartz, 20 to 50% plagioclase, 20 to 50% microcline, and 1 to 15% muscovite. The pegmatite consists of microcline, aplite, and biotite. The aplite dikes are cut by andesite and rhyolite dikes. A potassium-argon age determination—not considered reliable—gave an age of 30 (± 6) million years for this rock.

Andesite dikes intrude the granite, metamorphic rocks, and aplite-pegmatite dikes. Two andesite dikes cut volcanic rocks. Andesite dikes occur throughout the range but are much more abundant in the northern half. These dikes are up to 50 feet wide; most trend N 50° W in the north half of the range and northeast in the south half, and they dip steeply. The basic dikes range from diabasic to diorite in composition but mostly are hornblende



Scale: 1 inch = approximately 4 miles.

FIGURE 3. Geologic Map of the Sand Springs Range and Fairview Valley.
(Adapted from reference of footnote 2.)

andesites. Many of these dikes are porphyritic with phenocrysts of plagioclase and/or hornblende, while others are aphanitic. Some have porphyritic centers and aphanitic margins. The hornblende andesite contains 30 to 65% plagioclase (An_{30-55}), 15 to 60% hornblende, 2 to 5% magnetite, 1 to 5% biotite, 1 to 3% quartz, and less than 1% apatite. They are propylitically altered. Alteration minerals include chlorite, epidote, calcite, and quartz. In most cases, the andesite dikes are cut by the rhyolite dikes, but the reverse relationship has been observed (footnote 1).

Rhyolite dikes, up to 50 feet wide, intrude the granitic and metamorphic rock. They cut the aplite-pegmatite dikes and most of the andesite dikes. They trend N 50° W in the north half of the range, northeast in the south half, and have steep dips. Many are straight, long, and of even width, but the younger rhyolite dikes are curved, short, and variable in width. They range from rhyolite to dacite in composition. The three most common varieties are aphanitic rhyolite, quartz, porphyry with feldspar phenocrysts, and porphyry with both quartz and feldspar phenocrysts. The dikes are white to buff but commonly stain brown to reddish-brown on weathered surfaces. The aphanitic rhyolite consist of intergrown quartz and feldspar. The feldspar commonly is almost completely argillized and sericitized. The porphyritic rhyolite contains up to 30% phenocrysts of quartz, sanidine, and/or albite in a groundmass of the same minerals, plus 0.5% biotite. Some of the porphyritic rhyolite exhibits propylization, argillization, sericitization, and/or pyritization.

An intrusive breccia, consisting of inclusions, generally andesite and rarely granite, in a groundmass of pink and white aplite and porphyritic rhyolite, is exposed along the northeastern front of the Sand Springs Range. The inclusions range from gravel-sized, subrounded pieces to masses which have dimensions of hundreds of feet. Inclusions make up 20 to 50% of the rock.

Quaternary-Tertiary volcanic rocks unconformably overlie the metamorphic rocks at both ends of the Sands Springs Range and lie directly on the granite in several hills on Fourmile Flat to the west. An upper possibly Quaternary unit of basalt flowed west with various degrees of angular unconformity upon a possibly Tertiary unit consisting of an upper member of light-colored rhyolitic pyroclastics and a lower member of dark-colored andesite flows.

Of particular interest when considering geothermal potential are the rhyolites north of U.S. 50 and east of Sand Springs Pass. The description in the first Shoal report (footnote 1) is as follows:

"North of U.S. 50 and east of Sand Springs Pass several small erosional remnants of rhyolite vitrophyre are present near the top of the rhyolitic sequence. Probably all of these are at the same stratigraphic position and represent one large tabular, horizontal, shallow, glassy intrusive body or possibly an extrusive dome. The base and top of the vitrophyre are not present in any one exposure, but the body is probably less than 100 feet thick.

"The andesitic lower sequence rests with pronounced unconformity on the Mesozoic (?) metamorphic rocks in the northern end of the range. Locally, the rhyolitic sequence rests directly upon the metamorphics, indicating a profound local unconformity between the andesitic and rhyolitic sequences. In the southernmost part of the map area both the Quarternary (?) basalt and upper Tertiary (?) rhyolitic sequences are present, but the andesitic sequence is missing."

It would be highly desirable to obtain isotopic ages on the younger rhyolites.

Quaternary and Tertiary alluvial and eolian deposits occupy the valleys east and west of the Sand Springs Range. Fairview Valley, the site of Range Bravo 17, is characterized by playa deposits. On the basis of geophysical studies done for Project Shoal, the unconsolidated deposits are estimated to be 5,800 feet thick near the west margin of the valley a few miles south of U.S. Highway 50.

Folding is not a prominent feature in the Sand Springs Range. The metamorphic sequence at the south end of the range forms the western limb of a south plunging anticline, the eastern limb being the Fairview Range. Small-scale folding is uncommon in the Sand Springs Range except along the granitic-metamorphic rock contact where some drag folding of the metamorphics has taken place.

The Sand Springs Range is a north-south trending fault block uplifted along a series of high angle northeast and northwest trending faults. The first set trends about N 30° E and is accompanied by closely spaced, parallel-fracture cleavage. The N 50° W trending faults are accompanied by parallel joints. Most of the aplite-pegmatite, andesite, and rhyolite dikes are intruded along these faults and joints. Although many of these faults are narrow, some are wide and others are grouped together to form fault zones. Most contain gouge and brecciated wall rock. Iron staining, bleaching, and propylization is common in and along the faults. Primarily, the movement has been vertical, and has been from hundreds to thousands of feet, but it has been accompanied by some horizontal movement. These faults have been active from shortly after the intrusion of the granitic rocks until recent times.

A thrust fault cuts the metamorphic rocks at the north end of the range. It has a gently-north-dipping, undulating surface. The upper plate is garnetiferous, recrystallized limestone. The lower plate is metamorphic rocks. Offshoots of the granite body are exposed in contact with the thrust. Along the sole of the thrust the rock is bleached and stained. The direction of movement is not known.

A well-developed fracture cleavage is present throughout much of the granitic body and commonly parallels the northeast trending faults.

Deposits of tungsten occur in the metamorphic sequence at both the north and south ends of the range, in most cases in limestone at contact with the granite body. Scheelite and minor powellite occur, with garnet, cordierite, diopside, and calcite; and quartz veins containing scheelite cut the replacement bodies and surrounding rocks.

A system of east-west veins dipping steeply south, known as the Summit King System, occurs in a fault zone crossing the range just south of U.S. Highway 50. The veins cut metamorphic rocks. Tertiary volcanic rocks and andesite and rhyolite dikes occupy many of the individual fault strands. The veins are offset short distances by cross faults. At the surface, the veins have a braided pattern. The main southernmost veins dip 70 to 80 degrees south, and the subsidiary veins to the north dip 50 to 70 degrees south and intersect the main vein at shallow depths. The veins are 2 to 10 feet thick, contain fine-grained to vuggy quartz, angular fault breccia, and locally abundant calcite and pyrite. Ore minerals are native gold, cerargyrite (AgCl), and argentite (Ag_2S). The ratio of gold to silver is 1:40 in the upper level but decreases to 1:80 in the deeper workings. Underground workings reach depths of 450 feet. Several million dollars worth of silver and some gold have been produced from this vein system.

GEOPHYSICS

As noted previously, detailed studies of the Shoal site had been made (footnotes 1 and 2). Geophysical data from those studies were used extensively for this evaluation.

GRAVITY

A gravity survey was run over the Sand Springs Mountains and adjacent valleys for the Shoal Project (footnote 1). A complete Bouguer Anomaly Map was prepared. Data from this map are given as Figure 4. Supplemental data were taken from the Reno sheet, Bouguer Gravity Map of Nevada³. A complete Bouguer Anomaly Map indicates that terrain corrections were made. Gravity data indicate that the eastern frontal fault separating Fairview Valley and the Sand Springs Mountains is a complex system of northwest and northeast trending faults rather than a simple single fault. This is to be expected.

The maximum depth of valley fill in Fairview Valley is 5,800 feet thick in the center of the valley, about 6 miles south of U.S. Highway 50 (the south edge of the range). The maximum at U.S. Highway 50 is about 5,600 feet.

AEROMAGNETIC SURVEY

An aeromagnetic survey, done for the Shoal Project under the supervision of the Nevada Bureau of Mines and given as Figure 5, indicates the depth of unconsolidated sediments in Fairview Valley as 5,700 feet (footnote 1), a depth within 100 feet of that determined by gravity studies.

³Nevada Bureau of Mines. *Map 58. Bouguer Gravity Map of Nevada, 1977*. Reno, Nev., University of Nevada, 1977. (Publication UNCLASSIFIED.)

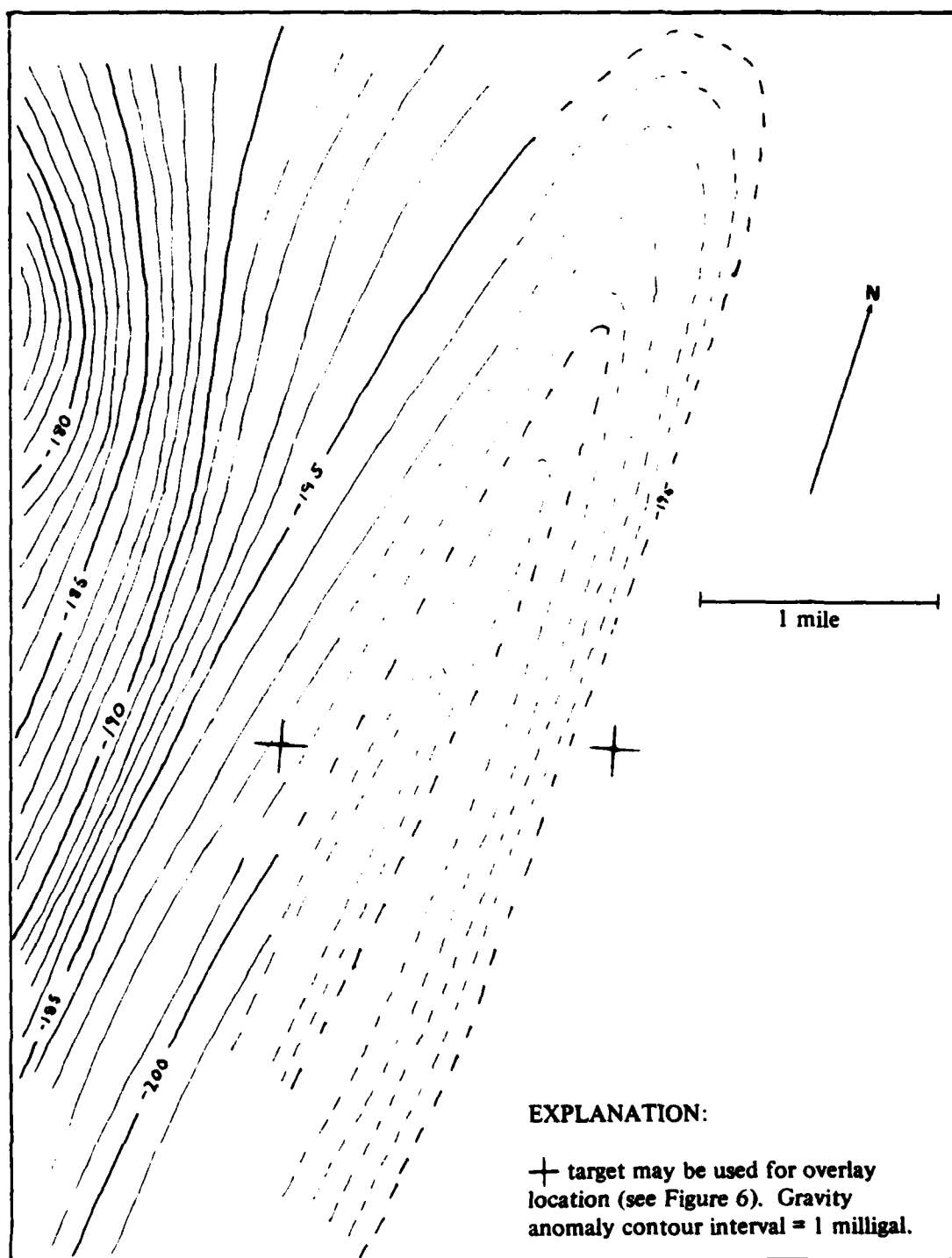
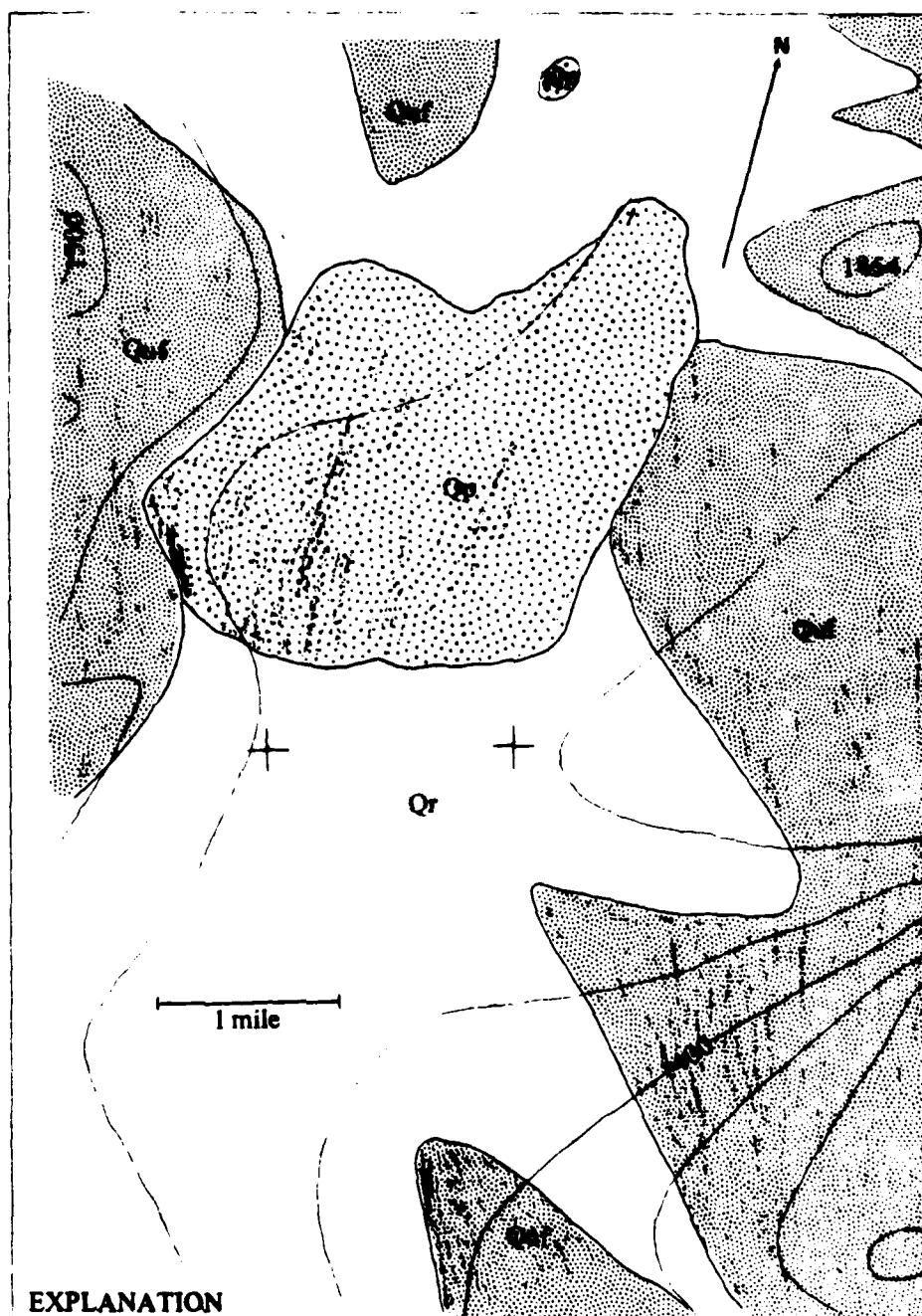


FIGURE 4. Gravity Map of the Sand Springs Mountains and Adjacent Valleys for the Shoal Project. (Solid contours adapted from reference of footnote 2. Dashed contours adapted from Bouguer Gravity Map of Nevada, Reno Sheet, Nevada Bureau of Mines Map 58, 1977.)



EXPLANATION

| | | | |
|-------------------|---|----|-----------------|
| Target | + | Qr | recent alluvium |
| Geologic contacts | — | Qf | fan deposits |
| | | Qp | playa deposits |

Magnetic contours show total intensity of magnetic field in gammas.

FIGURE 5. Aeromagnetic and Geologic Map of Target Area Range Bravo 17.
(Adapted from reference of footnote 2.)

SEISMIC REFRACTION STUDIES

A limited seismic study was made in the vicinity of Test Hole HS-1. This study indicated that the fill was compacted at about the elevation of the water table, approximately 310 feet below land surface datum (LSD), and that the depth to bedrock would be greater than 2,000 feet (footnote 1).

THERMAL GRADIENTS AND HEAT FLOW

Thermal gradient data were taken on three holes during Project Shoal. One hole, PM-1, was drilled to a depth of 1,340 feet in granite, near ground zero. The bottom hole temperature was 72°F. This translates to a thermal gradient of 3.0°C/100 m using a mean air temperature of 49.9°F. Normal gradients in the basin and range province are 3.0°C/100 m to 5.0°C/100 m. The curve showed no unusual characteristics (footnote 2, Plate 8).

Data were not available as to the length of time between hole completion and temperature logging. However, if the equilibrium gradient were high, one would expect the measured gradient to be higher than that obtained.

Also as part of the Project Shoal, thermal conductivity of the Sands Springs granite was determined. The four values obtained were 3.6, 3.1, 4.5, and 4.1 Cal/cm-sec-deg x 10⁻³ (footnote 1). The mean value was therefore 3.8 Cal/cm-sec-deg x 10⁻³. This is considerably lower than the values given by Birch and others⁴ for granite. The Shoal investigators attributed this to the granite having well-developed fracture cleavage.

Combining the thermal gradient and conductivity, one obtains a heat flow value of 1.15 Cal/cm²-sec (heat flow units (HFU)) which is slightly below average for the basin and range province (1.25 HFU).

Gradient data were also measured on the alluvium of Hole H-4 in Fairview Valley and Hole H-3 in Fourmile Flat. The bottom hole temperature of Hole H-3 was 58°F, which gives a gradient of 3.1°C/100 m. Hole H-4 had a bottom hole temperature of 69°F at 975 feet, which gives a gradient of 4°C/100 m.

Hole H-3 had a negative gradient from 329 to 400 feet. Hole H-4 had a negative gradient from 300 to 340 feet. These negative gradients may be caused by shallow ground water.

⁴Birch, Francis, J.F. Schairer, and H. Cecil Spicer. *Handbook of Physical Constants. Special Papers Number 36.* Boulder, Colo., Geological Society of America. January 1942 (reprinted 1961). 325 pp.

Thermal conductivity was not determined on the alluvium of Fairview Valley as part of Project Shoal. However, work done in 1978⁵ provided conductivity data for alluvium in the Monroe-Red Hill Hot Springs areas in Utah. Averaging these later data and converting to centimeter-gram-second(cgs) units would give $1.2 \text{ Cal/cm-sec-deg} \times 10^{-3}$ for dry alluvium and $3.76 \text{ Cal/cm-sec-deg}$ for saturated alluvium; therefore, the heat flow at Hole H-3 would be between 0.35 and 1.17 HFU and at Hole H-4 between 0.45 and 1.3 HFU. The higher values are similar to those obtained for the granite at the Shoal Site.

HYDROLOGY

Extensive hydrologic studies were made for the Shoal test. Data on wells and test holes used for the hydrologic studies are given in Tables 1 and 2. The two holes drilled in Fairview Valley are HS-1 and H-4.

Hole HS-1 was drilled in Quaternary alluvium consisting predominantly of sands with varying amounts of clay, silt, and gravel. Confined water was encountered in sand at a depth of 312.0 feet below the ground surface. Overlying the top of the sand aquifer was approximately 70.0 feet of fine-grained, silty sand. After the hole had been drilled to a depth of 530.0 feet below ground surface, a fine-grained, clayish, silty sand approximately 32.0 feet thick was encountered. This interval serves as an aquitard between the upper and lower aquifers, the latter consisting predominantly of sand with a higher percentage of silt and clay than the upper aquifer.

Test Hole H-4 was drilled in Quaternary alluvium consisting predominantly of sand with varying amounts of clay, silt, and gravel and Tertiary bedrock. Confined ground water was found in sand at a depth of 310.0 feet below ground surface. Overlying the top of the sand aquifer is approximately 65.0 feet of fine-grained, silty sand. After the hole had been drilled to a depth of 530.0 feet below ground surface, a fine-grained, clayish sand approximately 35.0 feet thick was encountered. This interval serves as an aquitard between the upper and lower aquifers.

At a depth of 813.0 feet below the ground surface, a distinct lithologic change was encountered with subsequent changes in the sample color, texture, percentage, and nature of the accessory minerals, along with an increase in drilling time required per foot. The possibility of Tertiary bedrock from 813.0 feet to the bottom of the hole at 935.0 feet exists. Hole H-4 was later plugged at a depth of 708 feet below the ground surface.

The well log of water point 18 (Table 3) was taken from the Sand Springs report (footnote 1). This well log is representative of sediments in the vicinity of Frenchman.

⁵U.S. Department of Energy, *Geophysical Study of the Monroe-Red Hill Geothermal System*, by W. Mose, D. S. Chapman, and S. H. Ward. Washington, D.C., DOE, 1978. (IDO/76-1601-77-17, publication UNCLASSIFIED.)

TABLE 1. Characteristics of Test Holes and Wells, Sands Springs Range and Fairview Valley.^a

| Designation | Location (U.S. Coast and Geodetic Survey coordinates) | Location (Federal land survey) | Elev. collar | Inclination | Vertical depth, ft | Diameter, in. | Use |
|----------------------|---|--|-----------------|-------------|--------------------------|-------------------------------------|--------------------------|
| Sands Springs Range | | | | | | | |
| EX-II-A ^b | N 1,619,292.72; E 558,740.32 | NE 1/4, NW 1/4, NW 1/4 Sec 3, T15N, R32E | 5158.90 | -----45" | 1342 | 2.9 | Shoal Proj. Test Hole |
| LC-II-D | N 1,619,975.66; E 557,545 | SW 1/4, SE 1/4; SW 1/4 Sec 34, T16N, R32E | 5238 | Vertical | 2017 | 6 | Shoal Proj. Test Hole |
| PM-1 | N 1,618,275; E 535,950 | SW 1/4, NW 1/4; NW 1/4 Sec 3, T15N, R32E | 5364 | Vertical | 1350 | 12.5 | Shoal Proj. Test Hole |
| PM-2 | N 1,622,985; E 558,200 | NF 1/4, SW 1/4; NW 1/4 Sec 34, T15N, R32E | 5325 | Vertical | 1310 | 12.5 | Shoal Proj. Test Hole |
| PM-3 | N 1,619,150; E 559,375 | NE 1/4, NW 1/4; NE 1/4 Sec 3, T15N, R32E | 5717 (?) | Vertical | 1110 | 12 (to 550 ft) 11.5 (to 1110 ft) | Shoal Proj. Test Hole |
| PM-8 | N 1,619,950; E 557,045 | SE 1/4, SW 1/4; SW 1/4 T16N, R32E | 5329 | Vertical | 900 | | Shoal Proj. Test Hole |
| USBM-1 | N 1,619,800; E 557,500 | SW 1/4, SE 1/4; SW 1/4 T16N, R32E | 5196 | Vertical | 1430 | | |
| Fairview Valley | | | | | | | |
| II-4 | N 1,622,285.67; E 576,914.39 | SW 1/4, SW 1/4; NW 1/4 Sec 32, T16N, R33E | 4244.92 | Vertical | 935 | 6 | Shoal Proj. Test Well |
| HS-1 | N 1,622,141.28; E 576,875.65 | SW 1/4, SW 1/4; NW 1/4 Sec 32, T16N, R33E | 4243.76 | Vertical | 699 | 8 | Shoal Proj. Test Well |
| WP-16 | | SE 1/4, SE 1/4; SW 1/4 Sec 32, T16N, R33E | 4262.78 | Vertical | 364 | 6 | Stock |
| WP-17 | | SW 1/4, SE 1/4; NW 1/4 Sec 3, T16N, R33E | 4153.3 | Vertical | 280 | 6 | Domestic |
| WP-18 | | SW 1/4, SE 1/2, NW 1/4 Sec 3, T16N, R33E | 4153.3 | Vertical | 288 | 8 | Domestic |
| WP-19 | | NF 1/4, NW 1/4; NE 1/4 Sec 11, T17N, R34E | 4147.80 | Vertical | 373 | 8 | Not Used (on Range) |
| WP-24 | | SW 1/4, Sec 18, T17N R34E | 4217 | Vertical | 334 | 6 | Stock |

^a From footnotes 1 and 2.^b Bearing - N 60° W, depth - 1898 ft.

TABLE 1. (Contd.)^a

| Designation | Discharge, gpm | Pump type | Depth to water, ft | Date | Temp. °f | Date |
|---------------------|-------------------|------------------------|-----------------------|---------|----------------------|---------|
| Sands Springs Range | | | | | | |
| FCH-A | | | Not encountered | 7/21/62 | See gradient data | |
| FCH-D | | | 968 | 8/1/62 | | |
| PM-1 | | | 1086 | 3/22/63 | | |
| | | | 1080 | 7/25/63 | | |
| PM-2 | | | 1100 | 4/24/63 | | |
| | | | 715 | 7/22/63 | | |
| PM-3 | | | 870 | 4/23/63 | | |
| | | | 1020 | 7/25/63 | | |
| PM-8 | | | 1070 | 7/2/63 | | |
| USBM-1 | | | 920 | 8/16/63 | | |
| | | | 925 | | | |
| Fairview Valley | | | | | | |
| H-4 | 70 | Submersible | 299.20 | 2/18/62 | 50 (at 320 ft) | 4/23/62 |
| IIS-1 | 70 | Turbine | 300.14 | 2/17/62 | 65 | 4/18/62 |
| WP-16 | 5 | Cylinder (Windmill) | 319.20 | 4/17/62 | | |
| | | | 319.28 | 7/10/62 | | |
| WP-17 | 15 | Submersible | 224.10 | 4/17/72 | 64 | 5/16/62 |
| WP-18 | 17 | Submersible | 224.60 | 4/17/62 | | |
| WP-19 | 5 | Cylinder | 218.84 | 7/10/62 | 60 | 5/22/62 |
| | | | 218.80 | 4/23/62 | 61 | 5/22/62 |
| WP-24 | 10 | Cylinder | 279.20 | 7/10/62 | | |
| | | | 299.70 | 5/9/62 | | |

^a from footnotes 1 and 2.

TABLE 2. Availability of Logs, and Drill Holes, Sand Springs Range and Fairview Valley^a

| Log | HOLE | | | | | | | |
|-----------------------------|-------|-------|------|------|------|------|--------|------|
| | ECH-A | ECH-D | PM-1 | PM-2 | PM-3 | PM-8 | USBM-1 | HS-1 |
| Core or bit size | X | X | X | X | X | | X | X |
| Casing | X | X | | | | | | |
| Water | X | X | X | | | | | |
| Percent core recovery | X | X | X | X | X | | | |
| Condition of core | X | X | X | | | | | |
| Drilling rate | X | X | X | X | X | | X | |
| Lithology | X | X | X | X | X | | X | X |
| Fractures and faults | X | X | X | X | X | | X | |
| Alteration & mineralization | X | X | X | X | X | | X | |
| Gamma ray | | | X | X | X | | X | |
| Neutron | | | X | X | X | | X | |
| Microlog | | | X | X | X | | X | |
| Electric | | | | | | | | |
| Spontaneous potential | | | X | X | X | | | |
| Resistivity | | | X | X | X | | | |
| Grain size-sludge | | | X | X | X | | X | |
| Caliper | | | X | | X | | X | |
| Gamma gamma density | | | X | X | X | | | |
| Sonic | | | | | | | | |
| Spontaneous potential | | | X | X | X | | | |
| Interval transit time | | | X | X | X | | | |
| Temperature | | | X | | | | X | |
| Fracture finder | | | | | | | X | |
| Micro seismogram | | | | | | | X | |

^aFrom footnotes 1 and 2.

TABLE 3. Well Log of Water Point 18.

| Depth | | Thickness, Feet | Type of Material |
|---------------|-------------|--------------------|---|
| From, Feet | To, Feet | | |
| 0 | 30 | 30 | Silt and Sand |
| 30 | 42 | 12 | Clay |
| 42 | 55 | 13 | Sand and gravel |
| 55 | 144 | 89 | Sand and gravel |
| 144 | 148 | 4 | Coarse gravel |
| 148 | 170 | 22 | White clay |
| 170 | 223 | 53 | Sandstone |
| 223 | 283 | 60 | Sand and silt with thin gravel layers—water encountered at 227 feet |
| 283 | 288 | 5 | White clay |

On 10 March 1962, Test Hole HS-1 was pumped for 20 hours, withdrawing from the interval between 310 feet to 530 feet, and Test Hole H-4 was used for recording observations (footnote 1).

On 5 May, the lower aquifer system (570 feet to 685 feet) at this site was tested using H-4 as a pump well and recording observations in Test Hole HS-1. For details of the test see the first Shoal report (footnote 1), pp. 56-60. Results are summarized in Table 4.

Movement of water in Fairview Valley is from south to north. From hydrologic data for Fairview Valley, the Nevada Bureau of Mines estimates the velocity of movement to be 25 feet or less per year.

Hydrologic data on the Sand Springs granite is harder to interpret. All holes drilled for the test event, except an inclined hole, ECH-A, encountered water.

The permeability of the granite is low (Table 4); however, fracture permeability occurs. A cross section of ground zero showing faults and shear zones is shown as Figure 3. Tests of Hole ECH-D (2,017 feet in depth) indicated that all fractures below a depth of 968 feet are saturated and yield water, but that fractures below 1,355 feet transmit the most water (footnote 1). Bailing tests of the various holes drilled in the granite zone gave rather inconclusive results. Part of the problem was that fractures contained water introduced during drilling. This added water has also complicated the interpretation of water chemistry. The transmissive constant, based on recovery curves after bailing, gave values for PM-1 of 28 to 77 gallons per day per foot (gpd/ft), PM-3 of 19 to 963 gpd/ft and USBM from 61 to 174 gpd/ft. Hole H-3, drilled in granite at Fourmile Flat, east of Sand Springs Range,

TABLE 4. Summary of Aquifer Characteristics, Sand Springs Range and Fairview Valley.^a

| Well and aquifer number | Specific capacity | Pumping test method | Coefficient of transmissibility (T), gpd/ft | Coefficient of storage (S) | Effective thickness, ft | Apparent permeability, gpd/ft |
|----------------------------|-------------------|---------------------|---|----------------------------|-------------------------|-------------------------------|
| H-4, HS-1 Upper Aquifer | 1.8 | Drawdown | 17,100 | 2.04×10^{-4} | 219 | 78 |
| H-4, HS-1 Lower Aquifer | 2.4 | Recovery | 11,000 | 2.45×10^{-4} | 120 | 91 |
| H-3 Bedrock | - | - | >200 | - | - | - |

| Permeability of Granite, millidarcies $\times 10^{-6}$ | Depth, ft |
|--|-----------|
| 13.8 | 512 |
| 3100 | 994 |
| less than 1 | 1422 |
| 390 | 2004 |

Core Drill Hole,

ECH-D

^aFrom footnote 2, pp. 75 and 59.

gave values greater than 100 gpd/ft. These values are, of course, low compared with those determined for alluvium on this project, 5,275 to 76,000 gpd/ft (footnote 2), but do indicate that the granite, where fractured, has appreciable permeability.

The Shoal shot raised the water level in Test Hole H-4, in Fairview Valley, 0.036 foot. The water returned to its preshot level within 24 hours. Water levels in holes within the granite rose as much as 65 feet and recovered to normal levels at variable rates.

GEOCHEMISTRY

WATER CHEMISTRY

Water chemistry and geothermometry values of water from six wells and test holes in Fairview Valley and five test holes in the Shoal Site (footnotes 1 and 2) are given in Tables 5 and 6. Unfortunately, the analyses are of poor quality. Of 21 analyses, six do not balance electrically within 20%. There are obvious differences between the same analyses given in parts per million (ppm) in footnote 1 and in equivalents per million in footnote 2. It is not possible to determine which is correct.

In Fairview Valley, the shallow aquifer waters are sodium-calcium-bicarbonate-sulfate waters. Potassium and magnesium are minor, and chloride is the least prevalent anion. The wells of the west side of the valley on the south end of the range produce water of good quality—260 to 282 ppm total dissolved solids (TDS). Well 24 on the west side of the Dixie Valley just north of the range has water of similar composition and quality (334 to 345 ppm TDS).

Water at Frenchman on the north central edge of the range is of poorer quality (664 to 1,014 ppm TDS) and may be characterized by potassium content exceeding calcium content. This could possibly indicate the leakage of geothermal waters into the aquifer. However, a more probable cause is precipitation on calcite as caliche due to evaporation from the playa with a resulting relative increase in potassium content. The waters of the playa dissolved solids from the alluvium as they moved from the edge of playa to the center and from the south to the north.

Chemical geothermometers exhibit great scatter (Table 5). Using only the analyses which electrically check within 20%, the mean silica reservoir equilibration temperature for the alluvial waters is 90°C and the Na-K-Ca temperature is 107°C. Using all of the available analyses, the silica temperature becomes 92°C and the Na-K-Ca temperature becomes 134°C.

TABLE 5. Composition of Water; Shoal Site, Sand Springs Range, and Fairview Valley.^a

| Item ^b | HS-1 | HS-1 | HS-1 | H-4 shallow aquifer | H-4 deep aquifer | PM-1 | PM-2 | PM-3 | ECH-D top sample | ECH-D middle sample |
|---------------------------------------|--------|----------|--------|---------------------|------------------|----------|----------|----------|------------------|---------------------|
| Date of sample | 4/4/62 | 11/26/63 | 2/1/64 | P5/28/62 | P7/11/62 | P5/29/63 | P5/29/63 | P5/29/63 | P7/19/62 | P7/19/62 |
| pH | 7.4 | 8.2 | 8.1 | 7.9 | 8.0 | 8.4 | 7.3 | 8.2 | 8.1 | 8.2 |
| TDS (ppm) | 282 | 262 | 283 | 265 | 338 | 1300 | 622 | 620 | 363 | 398 |
| SiO ₂ (ppm) | 38 | 58 | 67 | 51 | 59 | 19 | 3 | 54 | 24 | 13 |
| Equivalents per million | 2.30 | 2.26 | 2.18 | 2.64 | 2.98 | 2.26 | 6.62 | 1.40 | 3.96 | 2.36 |
| HCO ₃ as CaCO ₃ | ... | 0.02 | 0.12 | 0.02 | 0.20 | 0.02 | 0.12 | 0.04 | 0.16 | 0.16 |
| CO ₃ as CaCO ₃ | NA | NA | 0.04 | NA | NA | NA | NA | NA | NA | NA |
| NO ₃ ⁻ | 1.22 | 1.24 | 1.89 | 0.62 | 0.79 | 19.8 | 2.20 | 0.64 | 2.99 | 3.21 |
| Cl ⁻ | 0.89 | 1.12 | 1.06 | 0.81 | 0.87 | 2.70 | 1.62 | 1.27 | 1.77 | 2.04 |
| SO ₄ ²⁻ | 4.41 | 4.64 | 5.29 | 4.09 | 4.84 | 24.78 | 10.56 | 3.35 | 8.88 | 7.77 |
| Total anions | 3.35 | 2.00 | 2.00 | 4.70 | 2.17 | 11.81 | 3.05 | 17.80 | 3.53 | 3.21 |
| Na ⁺ | 0.85 | 0.18 | 0.18 | 0.23 | 0.12 | 0.23 | 0.23 | 0.46 | 0.13 | 0.10 |
| K ⁺ | 1.34 | 1.60 | 1.54 | 1.34 | 1.64 | 5.60 | 4.34 | 3.55 | 3.89 | 4.14 |
| Ca ²⁺ | 0.33 | 0.25 | 0.45 | 0.25 | 0.41 | 4.03 | 0.91 | 0.41 | 0.58 | 0.82 |
| Mg ²⁺ | 0.14 | 0.07 | ... | 0.11 | 0.14 | 0.04 | 0.04 | 0.18 | 0.04 | 0.04 |
| Fe ³⁺ | 6.01 | 4.10 | 4.17 | 6.63 | 4.48 | 21.71 | 8.57 | 22.50 | 8.17 | 8.31 |
| Total cations | PS-A | PS | PS | PS-A | PS | BS | BS-A | BS-A | PS | PS |
| Remarks ^c | 91 | 106 | 112 | 107 | 107 | 70 | 35 | 103 | 76 | 61 |
| SiO ₂ temp. ^{c,d} | 252 | 83 | 69 | 90 | 44 | 69 | 60 | 143 | 97 | 80 |

General note: NA = not analyzed; BS = bailed sample; PS = pumped sample; A = probable analytical errors.

^a From footnote 2, pp. 305 and 306.^b Hanks TDS at 600°C; data given in ppm; silica data in ppm; all other data in equivalents per million.^c All analyses by Hanks Laboratory except for HS-1, WP-18 (d), and WP-24, which were Hazleton-Nuclear Science Corporation, Los Angeles, California, and Las Vegas, Nevada (HNSC).^d Calculated by Geothermal Technology Branch, Public Works Department, Naval Weapons Center, China Lake, California.

TABLE 5. (Contd.)^a

| Item ^b | ICH-D Bottom Sample | ICH-D | USBM-1 | WP-16 | WP-17 | WP-18(a) | WP-18(b) | WP-18(c) | WP-18(d) | WP-24 | WP-24 |
|--|---------------------|-----------|--------|--------|--------|----------|-----------|----------|----------|----------|---------|
| Date of sample | P7/19/62 | 7/16/62 | 7/2/63 | 4/1/62 | 4/6/62 | P8/1/62 | P10/16/63 | 11/26/63 | 1/30/64 | P5/28/62 | 1/30/64 |
| pH | 8.1 | 8.5 | 8.0 | 7.4 | 7.6 | 8.3 | 8.5 | 8.6 | 8.4 | 8.2 | 7.8 |
| TDS, ppm | 400 | 388 | 463 | 260 | 918 | 990 | 1096 | 644 | 1014 | 334 | 345 |
| SiO ₂ , ppm | 15 | 49 | 31 | 50 | 50 | 70 | 40 | 14 | 65 | 29 | 35 |
| Equivalents per million | 2.40 | 3.68 | 2.36 | 2.32 | 7.82 | 9.50 | 8.36 | 8.78 | 8.44 | 2.24 | 2.22 |
| HCO ₃ ⁻ as CaCO ₃ | 0.10 | 0.30 | 0.02 | -- | 0.02 | 0.68 | 0.12 | 0.14 | 0.32 | 0.02 | -- |
| CO ₃ ²⁻ as CaCO ₃ | NA | NA | NA | NA | NA | NA | NA | NA | 0.01 | NA | 0.06 |
| NO ₃ ⁻ | 3.39 | 2.83 | 2.82 | 7.92 | 4.00 | 3.70 | 5.64 | 2.42 | 3.58 | 2.00 | 2.20 |
| Cl ⁻ | 1.88 | 1.91 | 3.35 | 0.88 | 3.88 | 3.84 | 8.12 | 3.73 | 3.92 | 1.21 | 1.21 |
| SO ₄ ²⁻ | 7.77 | 8.71 | 8.55 | 11.12 | 15.72 | 17.72 | 22.24 | 15.07 | 16.27 | 5.47 | 5.69 |
| Total anions | 3.35 | 3.74 | 4.53 | 3.30 | 20.65 | 15.00 | 17.75 | 10.73 | 14.80 | 6.05 | 4.52 |
| Na ⁺ | 0.21 | 0.23 | 0.10 | 0.64 | 0.95 | 0.46 | 0.26 | 0.26 | 0.26 | 0.15 | 0.16 |
| K ⁺ | 3.94 | 3.12 | 2.75 | 1.20 | 0.50 | 0.45 | 1.80 | 2.19 | 0.48 | 0.80 | 0.67 |
| Ca ²⁺ | 0.58 | 0.82 | 0.41 | 0.49 | 0.33 | 0.25 | 0.33 | 0.16 | 0.27 | 0.08 | 0.80 |
| Mg ²⁺ | 0.11 | 0.11 | 0.11 | 0.21 | 0.32 | 0.05 | 0.11 | 0.05 | 0.01 | 0.16 | -- |
| Fe ³⁺ | 8.19 | 8.02 | 7.90 | 5.84 | 22.75 | 16.21 | 20.25 | 13.39 | 15.82 | 7.24 | 6.15 |
| Total cations | PS | Air Lift | BS | PS-A | PS | PS | PS | PS | PS | PS-A | PS |
| Remarks ^c | | Sample HA | | | | | | | | | |
| SiO ₂ temp., °C ^d | 64 | 100 | 84 | 100 | 100 | 114 | 140 | 62 | 111 | 82 | 88 |
| Na-K-Ca Temp., °C ^d | 178 | 180 | 49 | 234 | 192 | 170 | 123 | 92 | 138 | 90 | 79 |

General note: NA = not analyzed; BS = bailed sample; PS = pumped sample; A = probable analytical errors.

^a From footnote 2, pp. 305 and 306.^b Hanks TDS at 600°C. data given in ppm; silica data in ppm; all other data in equivalents per million.^c All analyses by Hanks Laboratory except for HS-1, WP-18(d), and WP-24, which were HNSC.^d Calculated by Geothermal Technology Branch, Public Works Department, Naval Weapons Center, China Lake, California.

TABLE 6. Analyses of Waters From Drill Holes and Wells, Fairview Valley and Sand Springs Range.

| Item | IIS-1 | IIS-1 ^a | IIS-1 ^a | H-4 shallow aquifer | H-4 deep aquifer | PM-1 ^a | PM-2 ^a | PM-3 ^a | ECI-D top sample | ECI-D middle sample | ECI-D bottom sample |
|----------------------|-----------------|--------------------|--------------------|---------------------|------------------|-------------------|-------------------|-------------------|------------------|---------------------|---------------------|
| Date of sample | 4/4/62 | 1/26/63 | 2/1/64 | P5/28/62 | P7/11/62 | P5/29/63 | P5/29/63 | P5/29/63 | P7/19/62 | P7/19/62 | P7/19/62 |
| pH | 7.4 | 8.2 | 8.1 | 7.9 | 8.0 | 8.4 | 7.3 | 8.2 | 8.1 | 8.2 | 8.1 |
| TDS, ppm | 282 | 262 | 283 | 265 | 338 | 1300 | 622 | 620 | 363 | 398 | 402 |
| Ca | 27 | 32 | 44 | 27 | 33 | 112 | 87 | 71 | 78 | 83 | 79 |
| Mg | 4 | 3 | 5 | 3 | 5 | 49 | 11 | 5 | 7 | 10 | 7 |
| Na | 77 | 46 | 46 | 108 | 50 | 272 | 70 | 409 | 81 | 74 | 77 |
| K | 33 | 7 | 7 | 9 | 3 | 9 | 9 | 18 | 5 | 4 | 8 |
| Fe | 4 | 2 | -- | 3 | -1 | 4 | 1 | 5 | 1 | 1 | 3 |
| SiO ₂ | 58 ^b | 58 | 67 | 51 | 59 | 19 | 3 | 54 | 24 | 13 | 15 |
| SO ₄ | 53 ^c | 54 | 51 | 39 | 42 | 130 | 78 | 61 | 85 | 98 | 90 |
| Cl | 43 | 44 | 67 | 22 | 28 | 702 | 78 | 23 | 106 | 114 | 120 |
| CO ₃ | -- | 1 | 6 | 1 | 10 | 1 | 6 | 2 | 8 | 8 | 5 |
| as CaCO ₃ | | | | | | | | | | | |
| HCO ₃ | 115 | 113 | 109 | 132 | 149 | 113 | 331 | 70 | 198 | 168 | 170 |
| as CaCO ₃ | | | | | | | | | | | |

^a Calculated from equivalent part per million (footnote 2, pp. 305-306), by the Geothermal Utilization Division, Public Works Department, Naval Weapons Center, China Lake, California.

^b Value from footnote 1, p. 126. The value given in footnote 2, p. 305, is 38 ppm.

^c A value of 53 was reported in footnote 1, p. 126. A value of 43 was obtained by calculation of equivalents from footnote 2, p. 305.

TABLE 6. (Contd.)^a

| Item | EC/D Total Sample | USBM-1a | WP-16 | WP-17 | WP-18 | WP-18 ^a | WP-18 ^a | WP-18 ^a | WP-24 | WP-24 ^a |
|----------------------|-------------------|---------|--------|--------|--------|--------------------|--------------------|--------------------|----------|--------------------|
| Date of sample | 7/16/62 | 7/2/63 | 4/1/62 | 4/6/62 | 8/1/62 | P10/16/63 | 11/26/63 | 1/30/64 | PS/28/62 | 1/30/64 |
| pH | 8.5 | 8.0 | 7.4 | 7.6 | 8.3 | 8.5 | 8.6 | 8.4 | 8.2 | 7.8 |
| TDS, ppm | 338 | 463 | 260 | 918 | 990 | 1096 | 644 | 1014 | 334 | 345 |
| Ca | 63 | 55 | 24 | 10 | 9 | 36 | 44 | 10 | 16 | 13 |
| Mg | 10 | 5 | 6 | 4 | 3 | 4 | 2 | 3 | 1 | 10 |
| Na | 86 | 104 | 76 | 475 | 345 | 408 | 247 | 340 | 139 | 104 |
| K | 9 | 4 | 25 | 37 | 18 | 10 | 10 | 10 | 6 | 6 |
| Fe | 2 | 3 | 4 | 6 | 1 | 9 | 1 | 0.3 | 3 | -- |
| SiO ₂ | 49 | 31 | 50 | 50 | 70 | 40 | 14 | 65 | 29 | 35 |
| SO ₄ | 92 | 161 | 42 | 186 | 184 | 390 | 179 | 188 | 58 | 58 |
| Cl | 100 | 100 | 281 | 142 | 131 | 200 | 86 | 127 | 71 | 78 |
| CO ₃ | 15 | 1 | -- | 0.8 | 34 | 6 | 7 | 16 | 1 | -- |
| as CaCO ₃ | | | | | | | | | | |
| HCO ₃ | 184 | 118 | 116 | 391 | 475 | 418 | 439 | 422 | 112 | 111 |
| as CaCO ₃ | | | | | | | | | | |

^a Calculated from equivalent parts per million (footnote 2, pp. 305-306) by Geothermal Utilization Division, Public Works Department, Naval Weapons Center, China Lake, California.

The waters of the deep aquifer in Fairview Valley at Hole H-4 are chemically similar to those of the shallow aquifer with 338 ppm TDS. These waters give a silica temperature of 107°C and a Na-K-Ca temperature of 44°C.

Interpretation of the chemistry of waters from the Sand Springs granite is difficult. Water from HS-1 was used in drilling the test holes, and large quantities of water were lost. The holes were sampled almost immediately after drilling, so the waters were probably badly contaminated by drilling fluids. As noted earlier, HS-1 produced water of good quality with a low content of dissolved solids. The waters from the granite contained 363 to 1,300 ppm TDS. In general, the bicarbonate content exceeded the chloride content which exceeded the sulfate content. Hole PM-1 is an exception in that chloride exceeded bicarbonate. This could possibly indicate leakage from a geothermal reservoir, but the general poor quality of the analyses, coupled with the complete lack of other geothermal manifestations in the area, makes this interpretation doubtful. Sulfate is the predominate anion in the waters from Hole USBM-1, which could represent contamination from geothermal vapor; but again, considering the contaminated samples, the poor quality of analyses, and the lack of other indicators of geothermal potential, the writers do not feel such an interpretation is justified.

Silica temperatures of the analyses of waters from the Sand Springs granite, which balance electrically, range from 61°C to 100°C with a mean of 75°C. If all of the analyses are used, they range from 35°C to 103°C with a mean of 74°C.

Na-K-Ca temperatures of the better analyses range from 49°C to 180°C with a mean of 106°C. If all of the analyses are averaged, the mean is 107°C.

MERCURY GEOCHEMISTRY

Soil Sampling

As part of the geothermal exploration program, soil samples were collected on Range Bravo 17, on 7 February 1979, for subsequent mercury analyses. The sampling was performed by personnel of the Geothermal Utilization Division (John Bown, Jack Neffew, Ron Collins, and C. R. Rodgers) and assisted for safety reasons by naval Explosive Ordnance Disposal (EOD) personnel. Analyses of the samples collected are discussed in a later section of this report.

Sixty-three samples were collected along five traverses laid out along approximate north-south bearings. Traverse "A" started at the south end of what is known locally as the "airstrip" and extended north to U.S. Highway 50. Traverse "B" started at U.S. Highway 50 and ran south along the eastern flight line through the bullseye to the south boundary of the range. Traverse "C" ran along the centerline road from the north fence to the end of the road. Traverse "D" extended from the north fence south to the target along the

western flight line. Traverse "E" extended from the intersection of U.S. Highway 50 and Nevada 31 south along Nevada 31 for about 4 miles. The traverses were 0.5 to 1 mile apart. Samples were taken at 0.3 mile intervals using vehicle odometers. The traverses were partially staked.

On the advice of Mr. James Anderson, a geochemist with Pacific Energy and Minerals, Golden, Colorado, the top-most layer of soil was sampled. Loose gravel, snow, and organic matter were brushed away, and several ounces of soil were collected with plastic scoops and put into airtight plastic vials. Individual samples were labeled with a letter and a number.

Several problems hampered the sampling program at Range Bravo 17. No adequate map or recent aerial photograph seems to exist for this range. Because this is an active bombing range, certain areas could not be sampled for safety reasons. Also, only about one-third to one-half of the range was covered, since the eastern portion of the area reaches elevations over 8,000 feet and was not accessible due to a heavy snow cover. The final problems were also weather related. After a long period of freezing weather, a thaw had begun several days before sampling was performed. According to experts, a heavy snow melt or rainfall will lower the mercury content of soil. Also, the thaw caused several areas to become impassible due to mud.

Sample Preparation

The soil samples from Range Bravo 17 were damp to wet when collected; on return to the Naval Weapons Center, they were dried overnight at 125°F. After drying, the samples were screened through an 80-mesh stainless steel screen (180 micrometers, 0.007 inch) and stored in 76-cc plastic bottles prior to mercury determinations. Several samples, primarily the clays, had to be crushed prior to screening to provide sufficient sample for analysis. The samples were either (1) quartz sands or (2) silts and clays.

Mercury Analysis

Range Bravo 17 soil samples were analyzed for mercury content using Thermotron Corporation's atomic absorption mercometer in conjunction with a mercury collector unit. This unit collects the mercury vapor on silver wool, allowing extraneous gaseous materials to flow through the equipment prior to the mercury vapor count. The instrument is calibrated by injecting a known quantity of mercury vapor into the collector unit and noting the readout. For example, 1 cc of mercury vapor at 26°C contains 21.52 ng (10^{-9} gram) of mercury and might read 73 counts, each count being 21.52/73 or 0.295 ng per count.

In operation, the soil sample is weighed and placed in a quartz tube for heating. The gases, including the mercury vapor, are driven off the sample by heating between 650 and 799°C. (Nitrogen gas at 250 ml/min is used as the carrier fluid.) They then flow through water vapor and acid gas filters into the mercury collector unit where

the mercury vapor is absorbed on Collector A (the first of two silver wool collectors, A and C). Extraneous gases flow on through the instrument and are exhausted. Collector B is then heated, which drives the mercury vapor off and deposits it on Collector C. The mercrometer is then balanced to zero and the mercury vapors are driven off by heating Collector C. The imbalance between the detection cell and reference cell is displayed on the digital readout.

The Range Bravo 17 samples were run through twice. The first series was run using only the mercrometer, because the collector unit had been contaminated with a high mercury content sample. The tubes containing the silver wool were removed from the collector unit and baked out by the Chemistry Division at NWC, but since the series had been started with a particular setup, the collector unit was not put into operation until the first series was completed.

Duplicate, or additional, samples were run on both the first and second series. The first series produced mercury counts which were more tightly grouped than the second series, which was run using the collector unit. The instrument designer, Mr. Thomas Dooly (formerly with Thermotron Corporation), pointed out that calibrations made by directly injecting mercury vapor into the detection cell may be in error because the injection rate is unknown and the count is dependent on the percentage of mercury vapor in the cell. When the collector unit is used, however, the mercury vapor is driven off at a specific rate and is carried into the detection cell by a known and reasonably invariable quantity of air or nitrogen. The same reasoning applies to the sample if the gases from the sample (which are released at different temperatures) are led directly into the detection cell. For this reason, the use of the collector unit should result in more accurate calibrations and results.

Therefore, the second run was made using the collector unit in combination with the detector unit. About half-way through the samples, it was discovered that residual counts would show up if the collector unit was cycled without a sample or with the old sample still in the sample tube. Furthermore, the number of residual counts varied depending on the time selected for recycling. The run was completed recording both sample counts and periodic residual counts. The residual count varied slightly but usually was in the range from 6.5 to 11 counts with sample counts measuring from 6.8 to 11. Actually, the residuals could probably be ignored at these levels since they correspond to only 1.37 and 3.2 ng, respectively.

Computations

The quantity of mercury vaporized from each sample is calculated as nanograms of mercury per gram of sample. The calculation is made as follows: determine the quantity of mercury per calibration count and multiply this number by the number of counts obtained on running the sample to obtain nanograms mercury per gram of sample weight. In practice, the residual counts are subtracted from both the calibration and run counts before the calculation is made.

EXAMPLE:

Suppose the sample count is 28 for a 0.5 gram sample, the residual count is 8.0, and the calibration count is 62 with a mercury temperature of 20.0°C. Looking at the vapor-temperature table for mercury, 1 cc contains 13.19 ng which was injected into the mercury analyzer. Assuming that the counts are high by 8, subtract 8 from 28 and 62. Therefore:

$$\frac{13.19 \times (28 - 8)}{(62 - 8) \times 0.5} = 9.77$$

$$= 9.77 \times 10^{-9} \text{ grams Hg/gram sample}$$

Results

Results are given in Figure 6. The median value for mercury content is 8.85; the mean value is 9.7; the standard deviation is 4.1. Using two standard deviations as threshold value, the threshold would be 17.9. Only four samples exceeded the threshold and they produced one point anomalies.

Comparing the isoconcentration map with the geologic map, it appears that the deposits and younger alluvium are very low in mercury, while the alluvial fans coming off the mountains are somewhat higher in mercury content.

CONCLUSIONS

The geothermal potential of the Shoal area and Range Bravo 17 is considered low for the following reasons:

1. Thermal gradients in both areas are low.
2. Heat flow in both areas is low.
3. No hot or warm wells or springs occur in the areas, nor are there any fossil hot springs.
4. No hydrothermal alteration or mineralization of the type generally associated with geothermal resources was noted in the areas.
5. The silica and Na-K-Ca chemical geothermometers do not indicate that the waters have recently been in equilibrium with rocks at elevated temperatures.
6. The mercury content of the solid material of Fairview Valley is very low.
7. The valley fill in Fairview Valley is 5,800 feet deep. A geothermal resource, if present, would be very deep.

Because of the above-listed factors, further exploration at the Shoal Site and Fairview Valley is not warranted at this time.

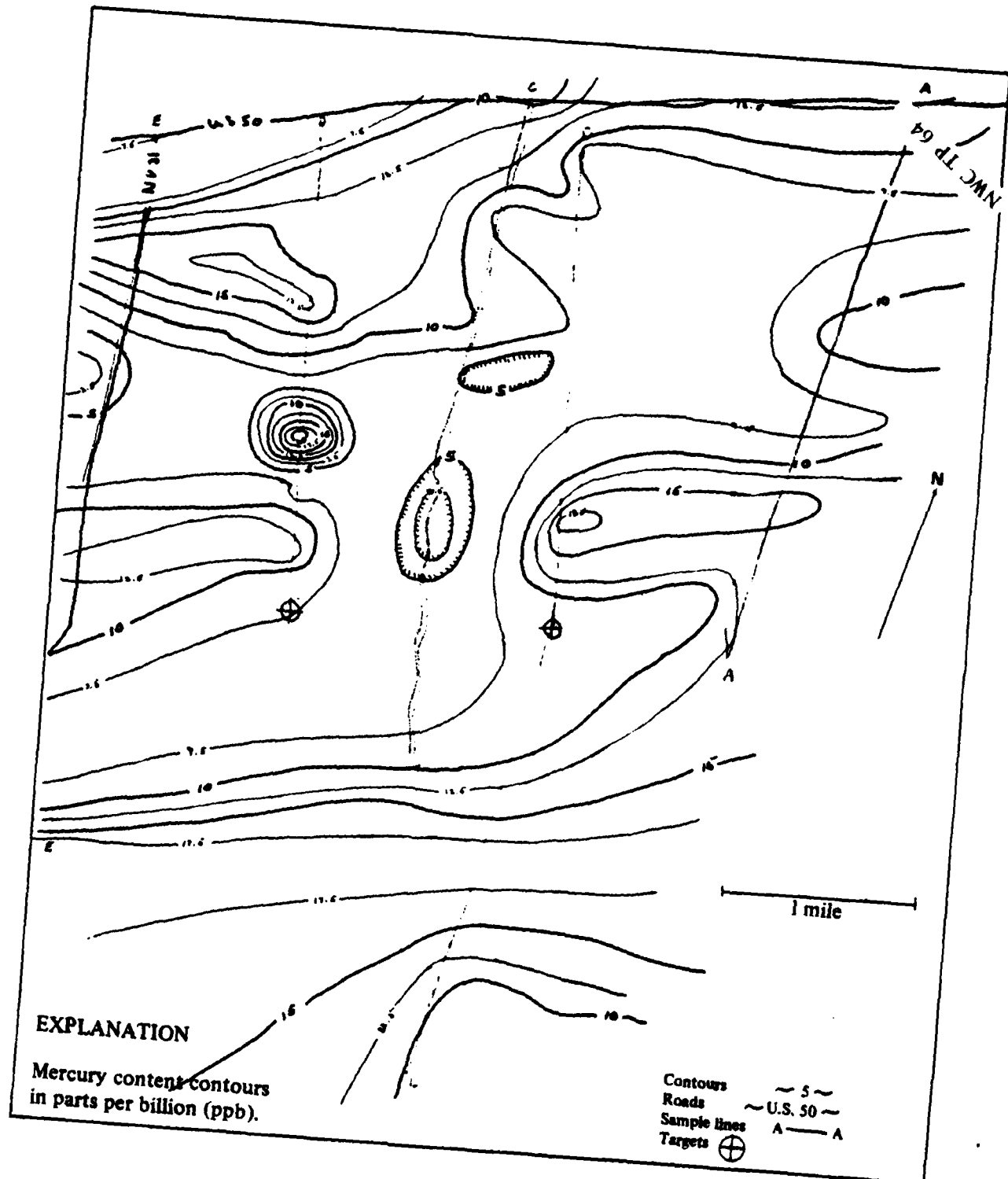


FIGURE 6. Mercury Content in Soils, Range Bravo 20, NAS Fallon.

Appendix A

STATION CLIMATIC SUMMARY

STATION CLIMATIC SUMMARY

PREPARED BY THE NAVAL WEATHER SERVICE ENVIRONMENTAL DETACHMENT, ASHEVILLE, N.C.
BY DIRECTION OF THE COMMANDER, NAVAL WEATHER SERVICE COMMAND



NARRATIVE SUMMARY

The Naval Air Station Fallon, Nevada is located 50 miles east of Reno. The terrain is largely sagebrush and desert. Numerous dry lakes dot the area. The mountains that surround Fallon act as a barrier to the influx of warm moist air from the Pacific Ocean. The Sierra range modifies approaching frontal systems so that they pass over Fallon with the moisture content and temperature markedly reduced.

November marks the beginning of the winter season which lasts until March. Cloudiness and cool temperatures predominate. Cold fronts often provide little more than gusty winds and scattered broken clouds. Zonal flow during the winter months produces fast moving low pressure systems with abundant moisture, resulting in severe turbulence and icing over the Sierra Nevada's.

Precipitation averages 4 inches during the winter months usually as light showers. Snowfall exceeding 2 inches is extremely rare. Comparatively warm minimum temperatures melt the snow in a day or two. Fog is an occasional problem in restricting visibilities, but rarely persists beyond the morning hours. December and Jan-

uary are the months of greatest fog frequency, occurring on an average of 4 days.

Winter storm tracks are generally located to the north of Fallon and spring like weather prevails. Occasionally a storm brings deteriorating conditions with fresh snowfalls and freezing temperatures.

May marks the beginning of summer which lasts until September. During the summer, hot dry weather dominates. Thunderstorms typically develop late in the afternoon in the foothills to the south. Strong gusty winds associated with these thunderstorms occasionally obscure visibilities with blowing dust. For the most part, these storms are confined to the surrounding foothills and rarely pass directly over the station. Dustdevils are quite frequent during the day over the alkali and sand flats. Only in July and August is the temperature consistently in the 90 degree range. Low humidity prevails through the summer which reduces the effects of high temperatures. Afternoon thunderstorms are primarily responsible for a summer average rainfall of 5 inches.

FALLON, NEVADA

TOTAL PRECIPITATION INCHES

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MONTHLY AND SEASONAL DEGREE DAYS

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| 1945-49 | 5 | 17 | 108 | 240 | 410 | 610 | 800 | 791 | 752 | 482 | 371 | 88 | 517 |
| 1950-54 | 5 | 17 | 108 | 240 | 410 | 610 | 800 | 791 | 752 | 482 | 371 | 88 | 517 |
| 1955-59 | 5 | 17 | 108 | 240 | 410 | 610 | 800 | 791 | 752 | 482 | 371 | 88 | 517 |
| 1960-64 | 5 | 17 | 108 | 240 | 410 | 610 | 800 | 791 | 752 | 482 | 371 | 88 | |

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